

Study of genesis of masteroon intrusive bodies(Northwest of Aligudarz)

Sayed Vahid Shahrokhi, **Mahdi Khodaparast, *Bahari, Fariba

*Geology group, Faculty of Science, Islamic Azad University , Khorramabad branch

**Geology group, Faculty of Science, Islamic Azad University , Meymeh branch

Abstract

In a region of Lorestan county (Northern Aligudarz) intrusive bodies with upper Cretaceous age have been formed; that altogether have northwest- southeast trend. The rocks of these bodies have composite within the range of granite to granodiorite and are intensely mylonitized by later tectonic events. These mass granitoides have intruded within a series of rock units which causes the formation of a contact metamorphic aureole in them. The oldest existent rock units included a thick sequence of alternated shale and sandstone of upper Triassic- Jurassic age, which have possessed a mild metamorphic. Therewith , the above thick rock sequence shows contact metamorphism by emplacement of intrusive bodies in some place , which its reflection is the formation of semi-schists, cordierite hornfels and schists. These rocks compositionally belong to the Calc-alkaline and based on computer-criteria, in orogenic belt, follow the tectonically regime of the VAG. The studies shows that the required vapor pressure for the formation of igneous bodies have been 1.5-2 kbar and the forming temperature sets between 700 to 750°C these mass granitoides are S- type granites and depth of their. Forming is upper 30 km.

Key word: Trend, Metamorphism, Calcalkaline, VAG, Tectonical regim , Granitoid

Introduction

Masteroon granitoid bodies with the coordination (coordinates latitude) of (33° 25' ,33° 33' N and 49° 27' , 49° 46' E)is located in 15 km, north of Aligudarz city and north east of Lorestan province. This mass is part of the accessory plain of internal Mesozoic in the Sanandaj-Sirjan zone from the perspective of plate tectonic subdivision (fig 1). These penetrating masses including some big and small ones, which are investigated in this paper. The igneous rocks generally have been injected shale and sandy formation in Jurassic area, and are belonged to the cycle which is shown by enclosing masses of Broujerd, and Hamedan. the composition of these stones are varied from the biotit-bearing granodiorite, with medium grain size to granite and they are porphyry in some regions. The tourmaline aplitic granite and Pegmatite in the contact of the intrusive Granodiorite in north, and west north of Aligoodarz are next levels of granite penetrating, and the Quartz veins which cross the Slates and Jurassic shale's is the last level of this penetrating (intrusive) cycle. The purpose of this paper is investigating the tectonomagmatic environment and penetrating mass forming situation, which one of them is called Masteroon granitoid body, in the west of Kondor village, and the other one which his called the Khorhe granitoide body, is in the first 5 km of the Aligoudarz–Khomein road and is crossed by this road. Here we call these various types by Masteroon granitoid bodies.

Discussion

In the west north part of Sanandaj-Sirjan, ie(Sanandaj,Sahnd,Hamedan ,Shahrekord ,and the region between Broujerd and Golpaigan, it is impossible to survey the laramian orogenic phase (Bellon and Broud,1975). The influence of this in the case study are some penetrating (intrusive) masses in shale and sandstone of jurassic, which had great impacts on them and usually convert them. This metamorphic complexes from penetrating masses toward the extreme layer are made of black hornfels, metamorphic sandstone, and sub mature micaceous schist with vein and veinlet of Silicic and cordierit and silimanit bearing micaschist and at last slate and directed shale, and a little part of submature schist (Shahrokhi,1381).

According to the results of radiometric age defining by the (Rb-Sr) method (Masoudi,1997), the mass replacing age of Arak of granitoid body which is located in the north part of Masteroon mountains, is 99 mil years and the smaller ones and pegmatite are about 52 mil years. Thus the replacing age of Masteroon granitoid bodies in west of Kondor village could be same with the last cretaceous time, and its differential-magmatic products are occurring simultaneous with the following Paleocene- Eocene era (Lotfi and Shahrokhi, 1382).

A microscopic view on the petrology of this penetrating (intrusive) mass shows the many various facies:1-Biotit and muscovite bearing granodiorite, which is the main stone of granitoid mass (pic 1) 2-Pegmatite Granodiorit, with a hypidiomorphic texture and mylonitic fabric (pic, 2) 3-Biyotite and muscovite mylonitic bearing aplite granitic as post magma differential with mineralized Siliceous vein and veinlet which are generated from Granite magma, under the influence of tectonic stress and got the mylonitic fabric characteristics (Shahrokh,1381)(pic 3).

The diagram (Cox, 1989), Alkali complex against SiO₂ (fig 2) and (Winchester and Floyd,1976)diagram (fig3) were used in chemical naming of the above samples.

Based on this, all the samples are placed in a range from granite to granodiorit. To determine the magma type which is forming the Granitoid mass, the triple (Irvin and Barragar, 1971) diagram of AFM was utilized. And all the samples were placed in the range of Calcoalkaline (fig 4). According to this, preface about magma type situation, and in order to determine the tectonomagmatic intrusive masses environment, the Rb cationic diagram was used against the Y +Nb (Pearce et al, 1989).All the samples were placed in the range of VAG or the volcanic arc district (fig 5). And the Rb (Condie, 1970) diagram was used against Sr. In order to determine the magma depth and according to that, the magma forming depth was determined higher than 30 km (fig 6).

In order to determine the granitoid mass forming temperature different methods have been used triple of Q-Qr-Ab diagram (Bowen, 1958) .The amount of each one of these was compared and estimated by the norms, and then they were transferred to the diagram. Based on this ,all the samples were placed in the temperature range of 700 C to 750 C (fig 7).And also it is possible to determine the transferred granitoid mass forming temperature using (Waily, 1978)diagram.

The presence of biotit as the only mafic mineral in this intrusive mass and crossing the Granite Solidos arc, with the temperature of 720 C and the pressure of 1.5 bar it is concluded that in the pressure of 1.5 to 2 bar biotit has been crystallized from the melted Granite with 720 C. (fig8)

And also using the ACF diagram (fig 8) it is possible to express the Granite types. All the samples are placed in the range of S type granite. Other instances Of chemical analysis and data are confirming the sedimentary origins of underlined Masteroon granitoid bodies (fig 9).

Conclusion and suggestions

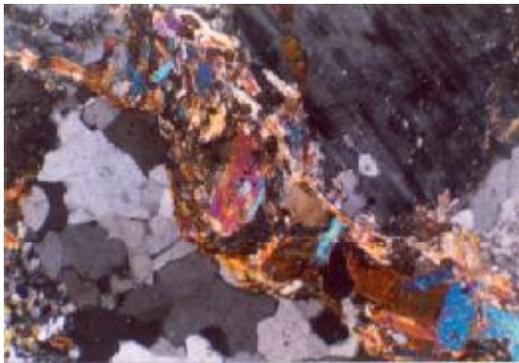
The replacing age of the Masteroon granitoid bodies is the late of the Cretaceous area to the early years of Paleocene which are related to Laramian orogenic.

The petrology composition of this mass is Granite to Grandiorite and its magma is of the Calco Alkaline, and belongs to the volcanic arc district and sedimentary origins. The mass forming temperature is 700°C to 750°C and the forming depth is more than 30 km .For more precise studies in field surveys the intrusive mass and type of area's fault and their numbers is investigated carefully. If possible the exact age of intrusive mass is determined by radio isotope methods.

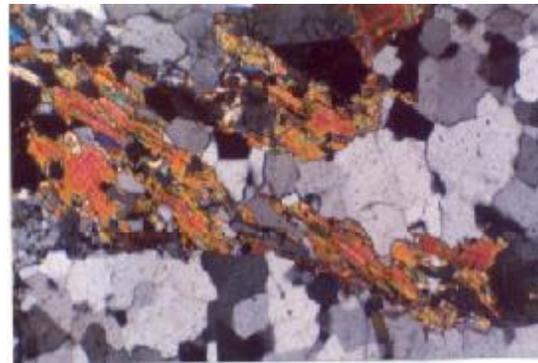
Refrence

- 1-Bagheri,N.,1380, The study of petrology, geochemistry and rare soils in granitoid bodies of north, northeast, northwest of Aligudarz with emphasis on anclave bodies , thesis of Esfahan university
- 2-Bellon, H., Braud,J., (Sans date): Donnes novells sur Le domaine metamorphique du Zagros (Zone de Sanandaj- Sirjan) au niveau de Kermanshah-Hamedan (Iran): Nature, age et interpretation des series metamorphiques et des intrusious, evolution structurale , fac, Sci. d'orsay, paris 14.p.
- 3-Condie, K.C., 1989. Geochemical changes in basalts and andesites across the Archean–Proterozoic boundary: identification and significance. *Lithos* 23, 1–18.
- 4-Cox, K. G., Bell, J.D., Pankhurst, R.J., 1979. The interpretation of igneous rocks. George Allen and Unwin., 450p.
- 5-Irvine, T.N., Baragar, W.R.A., 1971. A guide to the chemical classification of the common volcanic rocks. *Can. J. Earth Sci.* 8, 523-484.
- 6-Lotfi,M., 1381, The map of geology (1:100000) Mahneshan area with summary description, Geological survey of Iran
- 7-Maniar, P.D., Piccoli, P.M., 1989. Tectonic discrimination of granitoids. *Geological Society of America Bulletin* 101, 635–643.
- 8-Masoudi, F; 1997. Contact metamorphism and pegmatite development in the region SW of Arak, Iran, Unpublished PhD Thesis, Leeds University, UK.
- 9-Pearce, J.A., Harris, N.B.W., Tindle, A.G., 1984. Trace element discrimination diagrams for the tectonic interpretation of grantitic rocks. *Juornal of petrology* 25, 956 – 983.
- 10-Shahrokhi, S.V., Study of genetic & mineralization of Gholgholeh & Kervian gold deposite in Saghez area -Kordestan Province, Ph.D. Thesis, Tehran Science and Research Campus branch, Islamic Azad University, (2008)

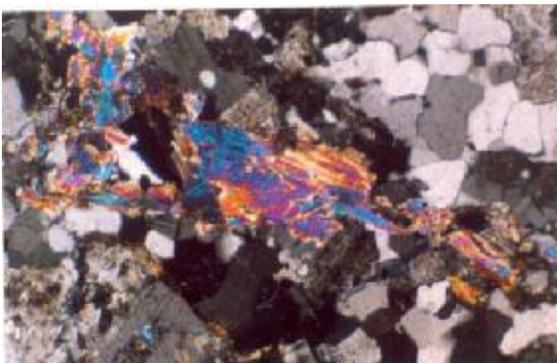
- 11-Shahrokhi,S., Lotfi,M., 1382, The study of control factors of Cu and Au mineralization and other elements related in Kondor region of Aligudarz, and it relationship with Geodynamic issues of Masteroon granitoid body ,(northeast of lorestan province), The seventh meeting of Geological society of Iran , Esfahan, Iran.
- 12-Soheyli,m., jafariyan,m., abdollahi,m., 1371, The map of geology (1:100000)Aligudarz area with summary description,Geological survey of iran.
- 13-Thiele, O., Seyed-Emami, K., 1968. Geological observation in the Borujird- Shahzand Area. Tehran, Geological survey of Iran, Mimeo, 4 Copies, 1 map (Inc copy 1 only).
- 14-Winchester, J.A., Floyd, P.A., 1975. Magma type and tectonic setting discrimination using immobile elements. Earth and Planetary Science Letters 27, 211–218.
- 15-Tuttle, F., Bowen, N.L., 1958, origin of granite in the light of experimental studies in the system NaAlSi₃O₈-KAlSi₃O₈-SiO₂-H₂O, geological science of America, 74, pp.153.



Picture1: View of the mylonitic fabric in granodiorite X(6.3*12.8)PPL



Picture2: view of the mylonitization in pegmatoid granodiorite in which mica series mylonitic fabric take control and that location is clear quartz grains intergrowth truly X (6/3*12/5)PPL.



Picture3: view of mylonitic fabric and series orientation of muscovite minerals in granitic aplite. In this picture quartz grain intergrowth, crystal of microcline, potassic feldspar, Albit are observable. X (6/3*12/5)PPL

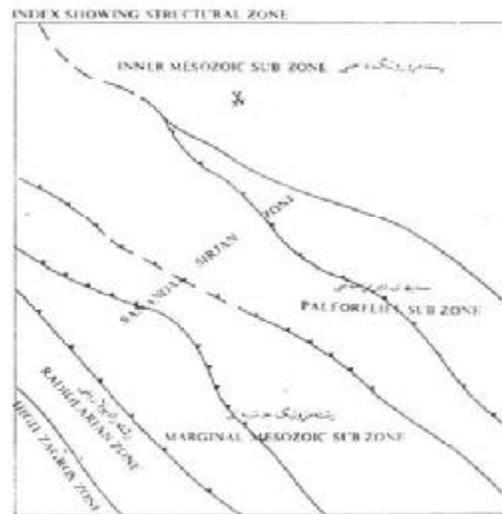


Figure 1: Aligoudarz divided region to sub-structural zone (Sohaili et al, 1371) and situation of Masteroon granitoid bodies.

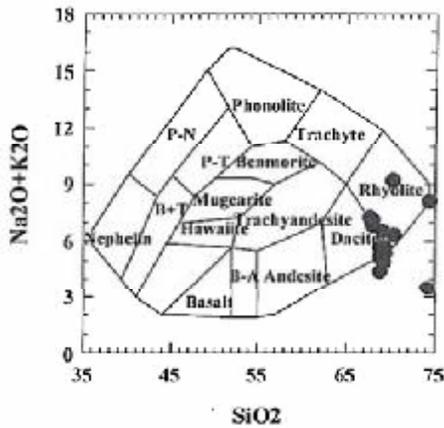


Figure 2: Cox diagram 1980

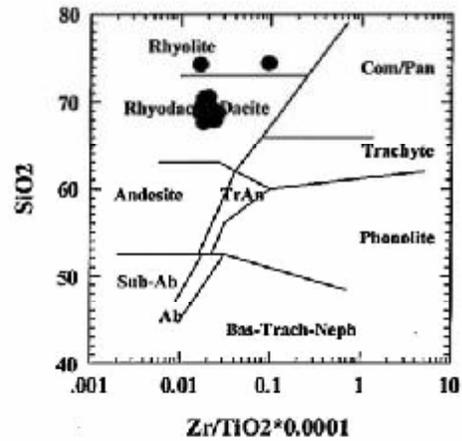


Figure 3: Winchester and Floyd diagram (1976)

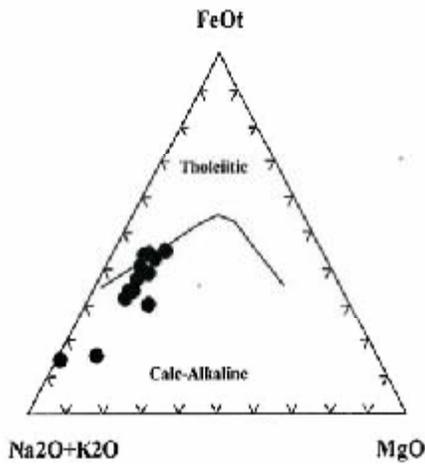


Figure 4: AFM diagram, (Irvin & Baragar, 1989)

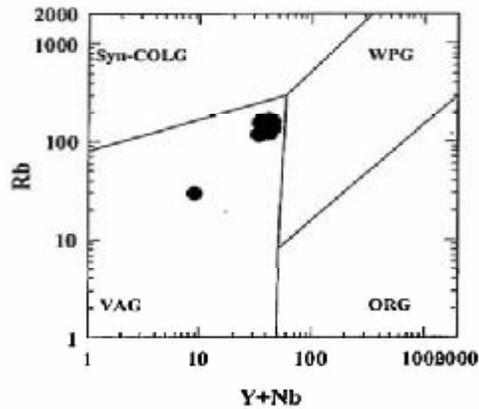


Figure 5: WPG unrognic granite, VAG volcanic arc district, ORG midocenic granit, continental plate collision (Syn-COLG)

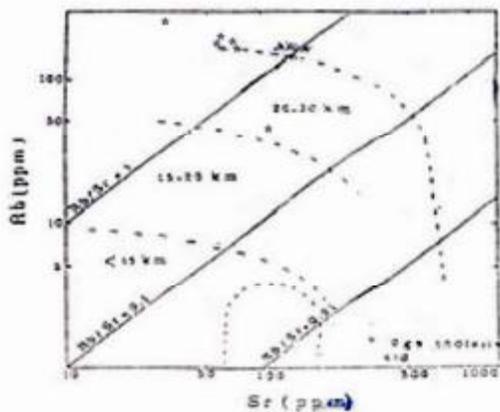


Figure 6: kondie diagram, 1970

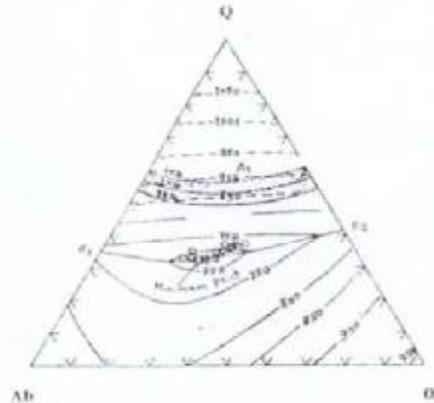


Figure 7: (Q-Ab-Or) system and situation of granitoid samples on it. M-minimum point of temperture in water saturated (Tuttle and Bowen, 1958) A- minimum of temperture in anhydrous state (Hings, 1985)

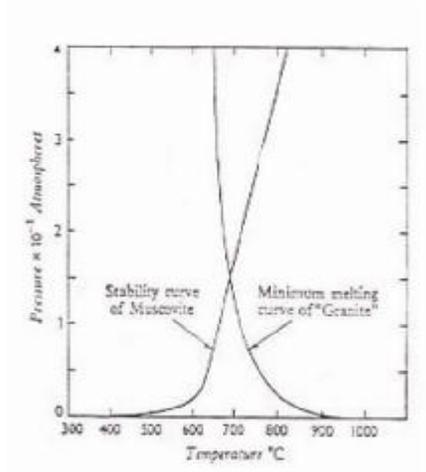


Figure 8: stability curve diagram of muscovite, biotite, hornblende, flogopite and melting curve of granite (viely, 1978)

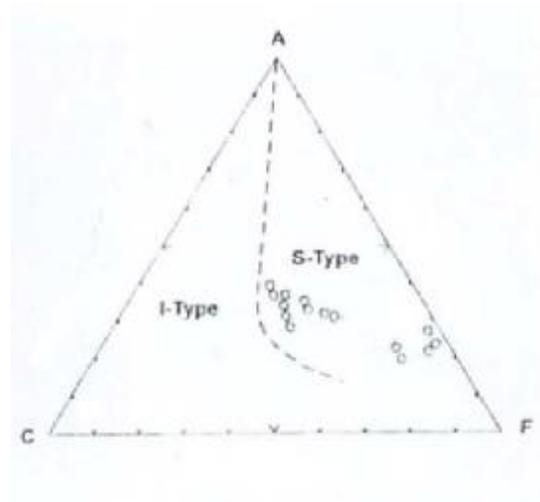


Figure 9: ACF diagram to determine type of granitoids of Aligudarz region. The data are as molly.
 F= FeO + MgO A= Al₂O₃ - Na₂O - K₂O
 C= CaO